

Norfolk Rodgers Bedrock Compilation Sheet 2 (paper)

Map

NOTICE !

Bedrock quadrangle 1:24,000 scale compilation sheets for the Bedrock Geological Map of Connecticut, John Rodgers, 1985, Connecticut Geological and Natural History Survey, Department of Environmental Protection, Hartford, Connecticut, in Cooperation with the U.S. Geological Survey, 1:125,000 scale, 2 sheets. [minimum 116 paper quad compilations with mylar overlays constituting the master file set for geologic lines and units compiled to the State map, some quads have multiple sheets depicting iterations of mapping]. Compilations drafted by Nancy Davis, Craig Dietsch, and Nat Gibbons under the direction of John Rodgers.

Geologic unit designation table translates earlier map unit nomenclature to the units ultimately used in the State publication.

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INTRODUCTION
Most of the unconsolidated (surficial) material in this quadrangle is till deposited by glacial ice that came from the north-northwest. The rest is stratified drift deposited by glacial melt water, postglacial alluvium, peat, and artificial fill. No weathered bedrock or other preglacial unconsolidated materials have been recognized.

Norris W. Hilary assisted with much of the field work and made most of the pebble counts. George P. Sweeney of the Soil Conservation Service, U.S. Department of Agriculture, kindly made available an unpublished soil map of Litchfield County from which the distribution of swamp and marsh deposits and areas of abundant bedrock exposures was in part derived. William M. Brown of the Soil Conservation Service provided subsurface data on the Norfolk Brook damsite and on proposed damsites on Wood Creek and Spaulding Brook. The Connecticut State Water Commission provided information on water wells.

TILL
Till deposited by ice that came from the north-northwest constitutes more than 90 percent of the unconsolidated material in the Norfolk quadrangle. Glacial grooves and strae at dozens of localities bear south-southeast (maximum range S. 2° E. to S. 53° E.). The mapped distribution of marble boulders in the northwestern Connecticut (Rodgers and others, 1959) implies that marble erratics scattered in the drift came from the northwest. The few till hills with glacial streamline forms (including drumlins) are elongate south-southeast. The till in these hills is of subglacial (lodgement) origin, whereas the till in the moraine-like ridges west of Norfolk village is probably ablation till.

In other parts of Connecticut Pessi and Schaefer (1968) have distinguished a widely distributed sandy, loose, relatively unoxidized upper till from a less widely exposed silty, compact, jointed, oxidized lower till that underlies it. The till in this quadrangle probably belongs to their upper till, though few good till exposures were seen and lower till is undoubtedly present in places.

Pessi and Schaefer (1968) believe that the upper part of their upper till may have been deposited by ice that came from the northeast. Pebbles derived from Triassic rocks of the Connecticut Valley found in a gravel pit half a mile east of the Norfolk quadrangle (Wangsten, 1970) give inconclusive support to this inference, and suggest that ice from the northeast may have entered at least the northeast part of this quadrangle.

EARLY ICE-CONTACT STRATIFIED DRIFT DEPOSITS
The ice front advanced progressively, though probably with fluctuations and with an irregular margin, from southeast to northwest across the area. Thus ice-contact stratified drift in the southeast (Qcd₁) is probably the oldest, having been deposited while ice covered most of the quadrangle. These deposits include rounded gravelly (kame) south of Winchester Center and a kame terrace northeast of the Park Pond dam; they may be approximately contemporaneous with ice-contact deposits mapped by Colton (1968) near Stillwater Pond, to the southwest in the adjoining West Torrington quadrangle.

After much of the ice at and near what is now Lake Winchester had melted, a melt-water stream drained southward across the Mad River-Naugatuck River divide south of Grant Swamp. Apparently this stream initially flowed in a tunnel under ice, where it formed an esker (part of Qcd₂ area). Later a kame terrace (part of Qcd₂) was deposited.

As the ice melted down, still further, a route was opened up for water to escape eastward down the Mad River valley. A group of kame terraces along Mad River that are apparently graded to the spillway east of Grant Swamp (Qcd₃) are lower and clearly younger than the kame terrace to the south (Qcd₂). Several small stratified and northeast of Dennis Hill are probably not contemporaneous with the deposits along Mad River, but are mapped as Qcd₃ because they are inferred from topography to be ice-contact origin, to be younger than the Qcd₂ kame terrace, and to have formed while ice still nearly filled the Blackberry River drainage basin.

Small patches of sandy drift on Ocean and Kugg Brooks cannot be reliably correlated with other deposits. They seem to be older than the ice-contact deposits along Mad River (Qcd₃), and have been designated Qcd₄.

OUTWASH IN HALL MEADOW AND RELATED DEPOSITS
A fan northeast of Tibbals Hill (Qohf) is probably contemporaneous with discontinuous terraces (Qoh) that extend down Hall Meadow and up Hall Meadow Brook. It is inferred that they consist of outwash from farther up Hall Meadow Brook because (a) I find no ice-contact features, yet (b) two lines of evidence suggest a melt-water origin (glacial rather than postglacial age). First, 4 of 100 pebbles collected for the gravel count at the pit southwest of Hall Meadow Cemetery were weathered that they crumbled in my hand, though associated pebbles of similar original lithology were fresh. Such fragile stonettes, presumably from preglacially weathered rocks, are not a normal constituent of Holocene deposits, but they are common in Pleistocene melt-water deposits; they were probably frozen solid when they were deposited. Second, angular erratic boulders scattered on the terrace deposits are reminiscent of the boulders, commonly ascribed to rafting on river-beds, that are common in many outwash deposits.

The tributary stream that enters Hall Meadow from the north follows the trough between the east-sloping Qohf fan and the valley wall. Clearly, aggradation of the fan has pushed it over to this position, and raised its profile. The stream from southeast of Parker Hill was only partially dammed, however, for no pond formed north of the fan, as commonly occurs in tributaries of an aggrading valley train. Instead, it built a bouldery, cobbly fan of its own (Qohp), which must have risen in step with the rising profile of the Qohf fan.

The deposition of the secondary fan (Qopf), in turn, pushed the stream from between Parker Hill and Turkey Cobble against the west wall of the valley. This stream, also, responded to the rising base level by depositing gravelly material (Qot), rather than by being dammed and forming a pond. Thus all three streams must have been heavily loaded and aggraded simultaneously. Both the secondary fan (Qopf) and the deposits to the north (Qot) are probably outwash deposits built by glacial melt water, though conceivably the streams that built them could have been fed by local rainfall.

The outwash and related deposits (Qohf, Qoh, Qopf, and Qot) are probably of about the same age as the ice-contact deposits south of Grant Swamp (Qcd₃). They are younger than the earliest ice-contact deposits (Qcd₁), for the ice had melted out from the valley at Hall Meadow before they were deposited, and they are probably older than the deposits along Mad River (Qcd₄), for ice probably filled the Mad River valley at the time they were deposited. Ice definitely filled this valley, standing at the divide northeast of Parker Hill, if the deposits north of Hall Meadow (Qopf and Qot) are outwash. In any case, ice probably filled the Mad River valley when ice standing near the site of South Norfolk contributed substantial volumes of heavily loaded melt water to build the Hall Meadow Brook deposits (Qohf and Qoh). The ice front is believed to have retreated in a generally northwesterly direction so that ice would have spilled across the saddle at the head of Mad River, north of Dennis Hill, whenever it spilled across the narrower saddle, more than 100 feet higher, at the head of Hall Meadow Brook, west of Dennis Hill.

LATER STRATIFIED DRIFT DEPOSITS
During deposition of the ice-contact deposits along Mad River (Qcd₄), ice almost filled the Blackberry River basin. Most stratified drift deposits in this basin are therefore younger than the Qcd₃.

Deposits east of Dutton Mountain.—In the wide lowland west and southwest of Dutton Mountain most of the stratified drift was deposited some distance behind the drainage divide. Much of the early deposition was over ice; Tobeys Pond is a large kettle, and the deposits to the south are now collapsed. Gravel in a road cut 1,900 feet west of corner 1293 is significantly coarser and thicker than the topset delta gravels exposed in the pit 700 feet to the northeast. The now-collapsed deposits were evidently thick enough to protect the ice they buried against melting for a time, so that their surface remained above the level of the glacial lake that formed later. They are grouped on the map with the deposits north of Spaulding Pond (Qcd₄), though they are probably not exactly contemporaneous.

After the ice surface farther north had melted down, a lake much larger than Spaulding Pond developed, held in by ice to the northwest. The water of this lake, glacial Lake Norfolk (Warren, 1969), escaped to Mad River by a spillway at about 1,325 feet elevation, 3,000 feet north of Dennis Hill. A large body of stratified drift, chiefly sand, accumulated in glacial Lake Norfolk as the Tobeys Pond delta (Qd). Much of the delta was built around and across remnants of ice, so that the delta surface is now extensively kilted and collapsed.

Deposits of stratified sand and silt (Qsu) in the upper Spaulding Brook valley, southeast of Tobeys Pond, are of uncertain origin and age. They contain scattered lenses of pebble gravel, and a few boulders as much as 4 feet in maximum diameter. Possibly they are collapsed ice-contact deposits contemporary with those to the northwest (Qcd₄), and owe their finer grain size to a position farther downstream where the gradient was reduced by the local base level of the spillway.

Lake deposits in Wood Creek valley.—The lake deposits (Ql) into which Wood Creek is incised consist of laminated silts with some layers of clay and very fine sand; an engineering report prepared for the Soil Conservation Service (W. M. Brown, oral commun., 1968) interprets them as varved. The same report also mentions soil-compression tests interpreted as showing that in the vicinity of the proposed dam the lake silts have been compacted by "sediments, since removed by erosion, which filled the valley to the level of the proposed dam" (indicated thickness of overburden more than 16 feet). I found no silts above about 1,195 feet elevation, though a well whose outlet is at 1,152 feet penetrated 68 feet of the lake silts. Lying on these lakebeds are scattered boulders and patches of till.

The scattered boulders and till suggest that ice probably covered the lake beds, though some boulders and till could perhaps have been rafted in on the lake by bergs, and some might have rolled or slid from silts above. Overriding by ice would also explain the results of the soil-compression tests: compaction by ice that subsequently melted seems a more likely explanation than compaction by strata entirely removed by stream erosion leaving remains.

The age relations of these lake deposits are uncertain. The lake beds are assumed in the map explanation to have been deposited in glacial Lake Norfolk. In this interpretation, if ice overrode the lake beds, it was a readvance postdating glacial Lake Norfolk, possibly the suspected ice movement from the northeast. It is not clear, however, why the silts should have such an apparently sharp upper limit near 1,195 feet if they were deposited in a lake whose surface was near 1,225 feet. The silts may have been deposited in a lake older than glacial Lake Norfolk, whose surface was near 1,200 feet and whose basin became almost filled by its own sediments. Any such low-level lake must have had an outlet down the Blackberry River. In this case, the ice readvance that closed the Blackberry River route and created Lake Norfolk presumably buried the Wood Creek lake deposits during the lifetime of Lake Norfolk. It is possible, however, that the Wood Creek lake beds are younger than glacial Lake Norfolk if the evidence for overriding has been misinterpreted.

Deposits west of Norfolk village.—Along the Blackberry River and south of Tobeys Pond Brook is a heterogeneous assemblage of materials ranging from clay and silt to cobbles, with patches of till. Laminated silt and clay with scattered boulders and inclusions of till were exposed by a cellar hole for a new house north of Sunset Drive; the bedding of the silt and clay was somewhat contorted. Neatly laminated sand and gravel, with only minor amounts of gravel, ridges of sand, gravel, and till at the head of the brook south of Case Pond have curved, moraine-like forms (Warren, 1969). Apparently all of this material was deposited in glacial Lake Norfolk, adjacent to the ice front as the front retreated northward, probably with fluctuations. Because the deposits lie northwest of the Tobeys Pond delta, they are assumed to be slightly younger. Though they are partly lacustrine, they are of ice-contact origin and are delineated as Qcd₅.

Glacial Lake Norfolk was finally drained when the ice melted sufficiently to permit flow down the Blackberry River valley, around the north and west sides of Cannon Mountain.

POSTGLACIAL DEPOSITS
Fans (Qf) are present at various places. Aggradation on some fans (Qfd) is apparently still in progress, for the depositing streams back up into distributary channels. The Norfolk Brook fan (Qfm) at Norfolk may have been built in late Pleistocene time, perhaps by mudflows (Warren, 1969). Terraces (Qtt) of sand and gravel, which stand above ordinary flood level, border streams at several places. Organic deposits (Qsm), in part thick peat, are present in many of the swamps and marshes. Alluvium (Qal) occurs along most streams, though commonly in belts too narrow to map.

ECONOMIC AND ENVIRONMENTAL CONSIDERATIONS
Materials of potential economic value in the surficial deposits of the Norfolk quadrangle include till, sand and gravel, water, and peat. Till suitable for use as fill is abundant throughout many parts of the quadrangle.

Economic resources of sand and gravel are generally scarce except in the area north and east of Tobeys Pond (Qd), where medium to fine sand, locally referred to as "dead" sand, is abundant and locally is more than 100 feet thick. Gravel is present south of Tobeys Pond, south of Grant Swamp, and in Hall Meadow. The gravels north of Spaulding Pond are of good quality, but any further pitting there might cause Spaulding Pond to drain into Norfolk Brook; this would adversely affect the Winsted city water supply, because these gravels provide water storage that maintains the level of Spaulding Pond and keeps Mad River perennial.

Properly screened and developed wells in the deltaic sands north of Tobeys Pond (Qd) would yield fair amounts of good water. Water is even more readily available from the gravels north of Spaulding Pond, but heavy pumping there would affect the discharge of Mad River.

Peat near Phelps Pond is of commercial quality for garden use as a soil conditioner (C. C. Cameron, oral commun., 1969). Peat is also present in Great Bear and Wildcat Swamps, in swamps south of North Pond and east of Dennis Hill, and in many smaller deposits.

The areas of thick till (Qt) symbol without bedrock exposures (overprint) are generally suited for disposal of solid waste by sanitary landfill operations. In areas of numerous outwash, unconsolidated material for borrow to bury the waste is scarce. Areas of stratified drift and of alluvium are poorly suited for sanitary landfill because of permeable soils and commonly because of a high water table. The delta sands (Qd) north of Tobeys Pond would be geologically suitable for disposal of fluid wastes, if the rate of addition was not excessive, as they would provide natural filtration to remove particulate solids and aeration to neutralize biologic oxygen demand; the water table is more than 50 feet deep in places.

ARTIFICIAL DEPOSITS
Artificial fill (Qat) is present in several places. It includes fill for a new house north of Sunset Drive; the bedding of the silt and clay was somewhat contorted. Neatly laminated sand and gravel, with only minor amounts of gravel, ridges of sand, gravel, and till at the head of the brook south of Case Pond have curved, moraine-like forms (Warren, 1969). Apparently all of this material was deposited in glacial Lake Norfolk, adjacent to the ice front as the front retreated northward, probably with fluctuations. Because the deposits lie northwest of the Tobeys Pond delta, they are assumed to be slightly younger. Though they are partly lacustrine, they are of ice-contact origin and are delineated as Qcd₅.

GLACIAL GROOVES, STRIAE, AND BOTH
Glacial grooves, striae, or both (G) are present in several places. They are generally oriented in a northerly to northeasterly direction, and are commonly associated with till. They are most abundant in the area north and east of Tobeys Pond (Qd), where they are generally oriented in a northerly to northeasterly direction, and are commonly associated with till. They are most abundant in the area north and east of Tobeys Pond (Qd), where they are generally oriented in a northerly to northeasterly direction, and are commonly associated with till.

MELT-WATER CHANNELS
Dry channel carved or modified by a glacial melt-water stream (generally limited in areas of abundant bedrock exposures). Terrace outlets (epitopes) of ice-margined lakes. Arrows indicate inferred direction of stream flow. Where channel is wide, banks are shown. Where one bank was ice, only the earth bank is shown.

HYPOTHETICAL SHORELINE OF GLACIAL LAKE NORFOLK
Based on the present topography, no shoreline features have been found except the spillway and delta (Qd). Bank marks (arrowheads) are shown only where they are believed to be reliable. No allowance made for postglacial tilt, which has affected the actual shoreline to an unknown degree.

DAMSITE
Damsite (D) shown by a rectangle with a vertical line through it. Centerline of proposed dam; some subsurface data available.

CONTACT, APPROXIMATELY LOCATED
Larger pits and areas of multiple pits are hatched to show approximate boundaries. Pits and quarries shown as abandoned probably have not been worked for more than a decade. Material symbols are identified below.

GLACIAL GROOVES, STRIAE, OR BOTH
Showing inferred direction of ice movement. Point of observation is at tip of arrow. Double-arrow symbol indicates a locality where grooves trending S. 20° E. are apparently cut by grooves trending S. 50° E.

LONG AXIS OF STRAHLINITE TILL DEPOSIT
Hill of till, with or without rock core, whose elongation is inferred to reflect the direction of ice movement (includes drumlins).

MELT-WATER CHANNEL
Dry channel carved or modified by a glacial melt-water stream (generally limited in areas of abundant bedrock exposures). Terrace outlets (epitopes) of ice-margined lakes. Arrows indicate inferred direction of stream flow. Where channel is wide, banks are shown. Where one bank was ice, only the earth bank is shown.

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EXPLANATION

- Alluvium**
Stratified sand, gravel, and silt. In many places, especially where swamps, undecomposed to partly decomposed organic material present at the surface. Along Mill Brook near east edge of quadrangle includes laminated silt, sand, and clay deposited partly in beaver ponds and probably also partly in a local Pleistocene lake. Holocene material at the surface probably overlies Pleistocene stratified drift in many places.
- Stream-terrace deposits**
Stratified sand, gravel (in places bouldery), and silt. Age uncertain, probably in places Pleistocene outwash, or ice-contact deposits without preserved bedrock, in places partly Pleistocene and partly Holocene and in places entirely Holocene.
- Fan deposits**
Gravel and sand. Qf age uncertain except that it postdates melting of ice at the site. Qfm, possibly mudflow fan of late Pleistocene age (Warren, 1969).
- Lake deposits**
Laminated silt, sand, and clay, possibly varved. In places overlies stratified sand and silt younger than glacial Lake Norfolk.
- Stratified drift?**
Stratified sand and silt with scattered lenses of pebble gravel and scattered angular boulders as much as 4 feet in diameter. Origin and correlation uncertain.
- Outwash(?) deposits**
Stratified sand and gravel. Qoh and Qohf are probably outwash. Qopf and Qot are contemporaneous but are possibly not of melt-water origin. Qohf, fan of Qoh with pebbles and silt. Qopf, sand to small boulders; fan deposited by stream from east of Parker Hill. Qot, sand to boulders, deposited by stream from east of Turkey Cobble.
- Till**
Nonstratified to poorly sorted mixture of sand and silt, clay, rock flour, and angular to rounded stones as much as 25 feet in maximum diameter; usually moderately compact. Does not occur with dilute boulders; where boulders, but commonly contains pockets of pebbly silt, sand, and silt. In areas lacking abundant bedrock exposures, till is discontinuous. Pockets or lenses of stratified sand and gravel commonly present. In areas of abundant bedrock exposures, till is discontinuous, commonly less than 10 feet thick, and includes many sandy areas in which undecomposed to partly decomposed organic matter is present near the surface. In areas lacking abundant bedrock exposures, thickness ranges from less than 1 foot (in near scattered outcrops) to a maximum reported figure of 75 feet. Most of the till is as young as the Qcd₃ deposits, but some till is as young as the Qcd₄ deposits, and possibly some till in the northeast dates from an ice advance postdating the Qcd₃.
- Artificial fill**
Artificial fill (Qat) is present in several places. It includes fill for a new house north of Sunset Drive; the bedding of the silt and clay was somewhat contorted. Neatly laminated sand and gravel, with only minor amounts of gravel, ridges of sand, gravel, and till at the head of the brook south of Case Pond have curved, moraine-like forms (Warren, 1969). Apparently all of this material was deposited in glacial Lake Norfolk, adjacent to the ice front as the front retreated northward, probably with fluctuations. Because the deposits lie northwest of the Tobeys Pond delta, they are assumed to be slightly younger. Though they are partly lacustrine, they are of ice-contact origin and are delineated as Qcd₅.
- Area of abundant bedrock exposures**
Pits and quarries shown as abandoned probably have not been worked for more than a decade. Material symbols are identified below.
- Glacial grooves, striae, or both**
Showing inferred direction of ice movement. Point of observation is at tip of arrow. Double-arrow symbol indicates a locality where grooves trending S. 20° E. are apparently cut by grooves trending S. 50° E.
- Long axis of strahlite till deposit**
Hill of till, with or without rock core, whose elongation is inferred to reflect the direction of ice movement (includes drumlins).
- Melt-water channel**
Dry channel carved or modified by a glacial melt-water stream (generally limited in areas of abundant bedrock exposures). Terrace outlets (epitopes) of ice-margined lakes. Arrows indicate inferred direction of stream flow. Where channel is wide, banks are shown. Where one bank was ice, only the earth bank is shown.
- Hypothetical shoreline of glacial Lake Norfolk**
Based on the present topography, no shoreline features have been found except the spillway and delta (Qd). Bank marks (arrowheads) are shown only where they are believed to be reliable. No allowance made for postglacial tilt, which has affected the actual shoreline to an unknown degree.
- Damsite**
Damsite (D) shown by a rectangle with a vertical line through it. Centerline of proposed dam; some subsurface data available.
- Contact, approximately located**
Larger pits and areas of multiple pits are hatched to show approximate boundaries. Pits and quarries shown as abandoned probably have not been worked for more than a decade. Material symbols are identified below.
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REFERENCES CITED
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SURFICIAL GEOLOGIC MAP OF THE NORFOLK QUADRANGLE, LITCHFIELD COUNTY, CONNECTICUT

By
Charles R. Warren
1972